

# Package: simET (via r-universe)

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**Type** Package

**Title** Tools for Simulation of Evapotranspiration of Field Crops and Soil Water Balance

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**Description** Supports the calculation of meteorological characteristics in evapotranspiration research and reference crop evapotranspiration, and offers three models to simulate crop evapotranspiration and soil water balance in the field, including single crop coefficient and dual crop coefficient, as well as the Shuttleworth-Wallace model. These calculations mainly refer to Allen et al.(1998, ISBN:92-5-104219-5), Teh (2006, ISBN:1-58-112-998-X), and Liu et al.(2006) [doi:10.1016/j.agwat.2006.01.018](https://doi.org/10.1016/j.agwat.2006.01.018).

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**Author** Minguo Liu [aut, cre], Huimin Yang [dct, fnd]

**Maintainer** Minguo Liu <liumg15@lzu.edu.cn>

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cal\_ActualVapourPressure\_for\_hourly

*Calculating actual vapour pressure for hourly time step*

### Description

Calculating actual vapour pressure for hourly time step

### Usage

cal\_ActualVapourPressure\_for\_hourly(Thr, RHhr)

### Arguments

Thr                    is average hourly temperature (degrees Celsius).  
 RHhr                  is average hourly relative humidity [%].

### Value

A vector for average hourly actual vapour pressure [kPa].

### References

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

cal\_ActualVapourPressure\_from\_dewPoint

*Actual vapour pressure derived from dewpoint temperature*

### Description

As the dewpoint temperature is the temperature to which the air needs to be cooled to make the air saturated, the actual vapour pressure is the saturation vapour pressure at the dewpoint temperature.

### Usage

cal\_ActualVapourPressure\_from\_dewPoint(Tdew)

**Arguments**

Tdew                dew point temperature(degrees Celsius).

**Value**

A vector for actual vapour pressure

**References**

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

---

cal\_ActualVapourPressure\_from\_psychrometricData  
*Actual vapour pressure (ea) derived from psychrometric data*

---

**Description**

The actual vapour pressure can be determined from the difference between the dry and wet bulb temperatures, the so-called wet bulb depression.

**Usage**

cal\_ActualVapourPressure\_from\_psychrometricData(Twet, Tdry, P, type)

**Arguments**

Twet, Tdry        wet bulb depression, with Tdry the dry bulb and Twet the wet bulb temperature (degrees Celsius).

P                is the atmospheric pressure (kPa).

type             psychrometer type ("Asmann type","natural ventilated","non-ventilated").

**Value**

A vector for Actual vapour pressure (ea)

**References**

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

---

cal\_ActualVapourPressure\_from\_RHmax

*Calculating actual vapour pressure derived from RHmax*

---

### Description

When using equipment where errors in estimating RHmin can be large, or when RH data integrity are in doubt, then one should use only RHmax.

### Usage

cal\_ActualVapourPressure\_from\_RHmax(Tmin, RHmax)

### Arguments

Tmin	daily minimum temperature (degrees Celsius).
RHmax	maximum relative humidity (%).

### Value

A vector for actual vapour pressure

### References

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

---

cal\_ActualVapourPressure\_from\_RHmaxAndRHmin

*Actual vapour pressure derived from RHmax and RHmin*

---

### Description

The actual vapour pressure can also be calculated from the relative humidity. Depending on the availability of the humidity data, different equations should be used.

### Usage

cal\_ActualVapourPressure\_from\_RHmaxAndRHmin(Tmax, Tmin, RHmax, RHmin)

### Arguments

Tmax	daily maximum temperature (kPa).
Tmin	daily minimum temperature (KPa).
RHmax	maximum relative humidity %.
RHmin	minimum relative humidity %.

### Details

For periods of a week, ten days or a month, RHmax and RHmin are obtained by dividing the sum of the daily values by the number of days in that period.

### Value

A vector for actual vapour pressure

### References

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

---

cal\_ActualVapourPressure\_from\_RHmean

*Calculating actual vapour pressure derived from RHmean*

---

### Description

In the absence of RH max and RHmin, it can be used to estimate actual vapour pressure.

### Usage

cal\_ActualVapourPressure\_from\_RHmean(RHmean, Tmax, Tmin)

### Arguments

RHmean	mean relative humidity(%).
Tmax	daily maximum temperature (degrees Celsius).
Tmin	daily minimum temperature (degrees Celsius).

### Value

A vector for actual vapour pressure

### References

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

---

cal\_afterRedistribution

*Calculating the volumetric water content after redistribution*

---

### Description

Calculating the volumetric water content after redistribution

### Usage

```
cal_afterRedistribution(THETA_v_sat, alpha, Ksat, deltaT, L, THETA11_1)
```

### Arguments

THETA_v_sat	soil saturation water content (m <sup>3</sup> m <sup>-3</sup> )
alpha	empirical coefficient. 13 for homogenous soil, 13-16 for heterogeneous soil
Ksat	saturated hydraulic conductivity
deltaT	time step difference (day)
L	the thickness(m) of soil layer i
THETA11_1	the volumetric water content before redistribution (m <sup>3</sup> m <sup>-3</sup> )

### Value

A value for the volumetric water content after redistribution(m<sup>3</sup> m<sup>-3</sup>)

---

cal\_airVaporPressureDeficit\_meanCanopyflow

*Calculating air vapor pressure deficit at the mean canopy*

---

### Description

Calculating air vapor pressure deficit at the mean canopy

### Usage

```
cal_airVaporPressureDeficit_meanCanopyflow(
  D,
  r_a_a,
  rho_cp = 1221.09,
  DELTA,
  A,
  gamma = 0.658,
  lambda_ET
)
```



**Arguments**

D	the vapor pressure deficit (mbar)
r_a_a	the aerodynamic resistance between the mean canopy flow and reference height (s m-1)
rho_cp	the volumetric heat capacity for air (1221.09 J m-3 K-1)
DELTA	is the slope of the saturated vapor pressure curve (mbar K-1)
A	is the total energy available to the system (W m-2 ground)
gamma	psychometric constant (0.658 mbar K-1 )
lambda_ET	the total latent heat flux (W m-2 ground)

**Value**

A vector for the vapor pressure deficit at the mean canopy flow (mbar)

**Note**

Knowing D0 is essential because this value is used to calculate the latent and sensible heat fluxes for the soil and crop components.

---

cal\_angerFromSouth      *Calculating anger from south*

---

**Description**

A parameter used to determine the position of the sun relative to the observer (the other one is solar inclination).

**Usage**

```
cal_angerFromSouth(latitude, solar_altitude, solar_declination)
```

**Arguments**

latitude	is the latitude data (Radian).
solar_altitude	It can be calculated from $\pi/2 - \text{cal\_solarinclination}$ .
solar_declination	is solar declination anger. It can be calculated from <code>cal_solardeclination()</code>

**Details**

The minus and positive signs are taken before and after solar noon, receptively. The reason for having the positive-and-negative signs is merely an artificial convention so that we are able to distinguish between the sun lying westwards (posite angles and after solar noon) and eastwards (negative angles and before solar noon).

**Value**

A vector for anger from south (Radian)

**References**

Teh CBS. Introduction to mathematical modeling of crop growth: How the equations are derived and assembled into a computer model. Brown Walker Press, 2006.

**Examples**

```
cal_angerFromSouth(latitude=0.52,solar_altitude=-0.715,solar_declination=-0.2974005)
```

---

```
cal_atmosphericPressure
```

*Calculating atmospheric pressure*

---

**Description**

The atmospheric pressure, P, is the pressure exerted by the weight of the earth's atmosphere.

**Usage**

```
cal_atmosphericPressure(elevation)
```

**Arguments**

elevation      elevation above sea level (m)

**Details**

Assuming 20°C for a standard atmosphere. Evaporation at high altitudes is promoted due to low atmospheric pressure as expressed in the psychrometric constant. The effect is, however, small and in the calculation procedures, the average value for a location is sufficient.

**Value**

A vector for atmospheric pressure (Kpa)

**References**

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

**Examples**

```
cal_atmosphericPressure(100)
```

---

cal\_bulkBoundaryLayerResistance  
*Calculating bulk boundary layer resistance*

---

**Description**

Calculating bulk boundary layer resistance

**Usage**

cal\_bulkBoundaryLayerResistance(nu, u\_h, w, L)

**Arguments**

nu	the wind speed extinction coefficient (taken as 2)
u_h	the wind speed at the canopy top (i.e., at plant height h) (m s <sup>-1</sup> )
w	is the mean leaf width (m)
L	leaf area index

**Value**

A vector for bulk boundary layer resistance (s/m)

---

cal\_canopyPenetrationProbabilityForNetRadiation  
*The canopy penetration probability for net radiation*

---

**Description**

The canopy penetration probability for net radiation

**Usage**

cal\_canopyPenetrationProbabilityForNetRadiation(kRn, L)

**Arguments**

kRn	the canopy extinction coefficient for net radiation (taken as 0.3)
L	the leaf area index (m <sup>2</sup> leaf m <sup>-2</sup> ground)

**Value**

The vector for canopy penetration probability for net radiation

---

cal\_canopyResistance    *Calculating canopy resistance*

---

**Description**

Calculating canopy resistance

**Usage**

cal\_canopyResistance(a1, a2, It, L, Lmax)

**Arguments**

a1, a2	are empirical coefficients, dependent on the crop type.
It	the total hourly solar irradiance (W m <sup>-2</sup> ground)
L	the leaf area index (m <sup>2</sup> leaf m <sup>-2</sup> ground)
Lmax	the maximum total leaf area index (m <sup>2</sup> leaf m <sup>-2</sup> ground)

**Value**

A vector for canopy resistance(s/m)

---

cal\_canopyTem    *Calculating canopy temperature*

---

**Description**

Calculating canopy temperature

**Usage**

cal\_canopyTem(Hc, r\_c\_a, Hs, r\_a\_a, rho\_cp, Tr)

**Arguments**

Hc	crop sensible heat fluxes (W m <sup>-2</sup> )
r_c_a	the bulk boundary layer resistance (s m <sup>-1</sup> )
Hs	soil sensible heat fluxes (W m <sup>-2</sup> )
r_a_a	the aerodynamic resistance between the mean canopy flow and reference level (s m <sup>-1</sup> )
rho_cp	the volumetric heat capacity for air (1221.09 J m <sup>-3</sup> K <sup>-1</sup> )
Tr	Tr is the air temperature at reference level (Celsius degree). weather station.

**Value**

A vector for the canopy (foliage) temperature (Celsius degree)

---

cal\_capillaryRise      *Calculating capillary rise*

---

### Description

Calculating capillary rise

### Usage

```
cal_capillaryRise(
  a1,
  b1 = -0.17,
  a2,
  b2 = -0.27,
  a3 = -1.3,
  b3,
  a4,
  b4,
  Dw,
  Wa,
  LAI,
  ETm
)
```

### Arguments

a1	Soil water storage to maximum root depth at field capacity(mm).
b1	A parameter.
a2	storage above the average between those at field capacity and the wilting point(mm).
b2	A parameter.
a3	A parameter.
b3	A parameter.6.7 for clay and silty clay loam soils, decreasing to 6.2 for loamy sands
a4	A parameter.4.6 for silty loam and silty clay loam soils,decreasing to 6.2 for loamy sands.
b4	A parameter.-0.65 for silty loam soils and decreasing to -2.5 for loamy sand soils.
Dw	Groudwater depth below root zone(m).
Wa	actual soil water storage in the root zone.
LAI	Leaf area index.
ETm	potential crop evaporanspiration (mm/day),usually ETm=ETc(mm/d).

### Value

The value for capillary Rise (mm/day).

**References**

Liu Y, Pereira L S, Fernando R M. Fluxes through the bottom boundary of the root zone in silty soils: Parametric approaches to estimate groundwater contribution and percolation[J]. Agricultural Water Management, 2006, 84(1):27-40.

---

cal\_cropRoughnessLength

*Calculating the crop roughness length*

---

**Description**

Calculating the crop roughness length

**Usage**

cal\_cropRoughnessLength(h)

**Arguments**

h                      the plant height (m)

**Value**

A vector for the crop roughness length(m)

---

cal\_daylightHours

*Calculating Daylight hours*

---

**Description**

Calculating Daylight hours

**Usage**

cal\_daylightHours(sunsetHourAngle)

**Arguments**

sunsetHourAngle

is the sunset hour angle in radians from cal\_sunsetHourAngle().

**Value**

A vector for day light Hours

## References

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

---

cal\_DeepPercolation     *Calculating Deep percolation*

---

## Description

Calculating Deep percolation

## Usage

cal\_DeepPercolation(Wa, Wfc, a, b, t)

## Arguments

Wa	actual soil water storage in the root zone (mm)
Wfc	soil water storage to maximum root depth ( $Z_r$ ) at field capacity (mm)
a	A water storage value comprised between WFc and Wa at saturation.
b	$b < -0.0173$ for soils draining quickly. Otherwise $b > -0.0173$ .
t	time after an irrigation or rain that produced a storage above field capacity (days)

## Value

A vector for deep percolation(mm/day).

## References

Liu Y, Pereira L S, Fernando R M. Fluxes through the bottom boundary of the root zone in silty soils: Parametric approaches to estimate groundwater contribution and percolation[J]. Agricultural Water Management, 2006, 84(1):27-40.

---

cal\_Dei\_for\_DualKc     *calculating the depletion in the topsoil layer at the end of the day*

---

### Description

In fact, it performs water balance in a day

### Usage

cal\_Dei\_for\_DualKc(Dei\_start, P, I, E, Dep, TEW)

### Arguments

Dei_start	Depletion in the topsoil layer
P	Precipitation
I	Irrigation
E	Evaporation on day i, mm
Dep	Deep percolation loss from the topsoil layer on day i if soil water content exceeds field capacity, mm
TEW	Maximum cumulative depth of evaporation (depletion) from the topsoil layer

### Value

A value for the depletion in the topsoil layer at the end of the day

---

cal\_DPe\_for\_DualKc     *Deep percolation loss from the topsoil layer*

---

### Description

Deep percolation loss from the topsoil layer

### Usage

cal\_DPe\_for\_DualKc(P, I, Dei\_start, fw)

### Arguments

P	Precipitation
I	Irrigation
Dei_start	Depletion in the topsoil layer
fw	Fraction of soil surface wetted by irrigation, 0.01-1



**Value**

A value for deep percolation loss from the topsoil layer

**References**

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

---

cal\_DPr\_for\_DualKc      *Deep percolation loss from the root layer*

---

**Description**

Deep percolation loss from the root layer

**Usage**

cal\_DPr\_for\_DualKc(P, Irrigation, ETa, Dri\_start)

**Arguments**

P	Precipitation
Irrigation	Irrigation
ETa	Actual evapotranspiration
Dri_start	Depletion in the root layer

**Value**

A value for deep percolation loss from the root layer

**References**

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

cal\_DP\_for\_singleKc     *Calculating deep percolation*

---

**Description**

Calculating deep percolation

**Usage**

cal\_DP\_for\_singleKc(P, I, ETa, Dri\_start)

**Arguments**

P	Precipitation
I	Irrigation
ETa	Actual evapotranspiration
Dri_start	The depletion of root layer

**Value**

A value for deep percolation

---

cal\_eddyDiffusivity\_Canopytop  
*Calculating eddy diffusivity at the canopy top*

---

**Description**

Calculating eddy diffusivity at the canopy top

**Usage**

cal\_eddyDiffusivity\_Canopytop(k = 0.4, u\_, h)

**Arguments**

k	the von Karman constant (0.4)
u_	the friction velocity (m/s)
h	the plant height (m)

**Value**

A vector for eddy diffusivity at the canopy top(m<sup>2</sup>/s)

---

cal\_eddyDiffusivity\_heightZ  
*Calculating eddy diffusivity at height z*

---

**Description**

Calculating eddy diffusivity at height z

**Usage**

cal\_eddyDiffusivity\_heightZ(Kh, nK, z, h)

**Arguments**

Kh	eddy diffusivity at the canopy top(m2/s)
nK	the eddy diffusivity extinction coefficient (taken as 2)
z	height(m)
h	the plant height(m)

**Value**

A vector for eddy diffusivity at height z (m2/s)

---

cal\_ET0\_from\_PM      *calculating reference evapotranspiration from Penman-Monteith method*

---

**Description**

The FAO Penman-Monteith method is maintained as the sole standard method for the computation of ETo from meteorological data.

**Usage**

cal\_ET0\_from\_PM(delta, Rn, G, gamma, Tem, u2, es, ea)

**Arguments**

delta	slope vapour pressure curve (kPa &deg;C). From cal_slopeOfSaturationVapourPressureCurve()
Rn	net Radiation at the crop surface [MJ m-2 day-1]. From cal_netRadiation()
G	soil heat flux density [MJ m-2 day-1].
gamma	psychrometric constant (kPa &deg;C).
Tem	air temperature at 2 m height [&deg;C].
u2	wind speed at 2 m height [m s-1].
es	saturation vapour pressure [kPa].
ea	actual vapour pressure [kPa].

**Value**

A vector for reference evapotranspiration [mm day<sup>-1</sup>].

**Note**

Ten-day or monthly time step :

Notwithstanding the non-linearity in the Penman-Monteith equation and some weather parameter methods, mean ten-day or monthly weather data can be used to compute the mean ten-day or monthly values for the reference evapotranspiration. The value of the reference evapotranspiration calculated with mean monthly weather data is indeed very similar to the average of the daily ETo values calculated with daily average weather data for that month.

When the soil is warming (spring) or cooling (autumn), the soil heat flux (G) for monthly periods may become significant relative to the mean monthly Rn. In these cases G cannot be ignored and its value should be determined from the mean monthly air temperatures of the previous and next month.

Daily time step:

Calculation of ETo with the Penman-Monteith equation on 24-hour time scales will generally provide accurate results.

As the magnitude of daily soil heat flux (G) beneath the reference grass surface is relatively small, it may be ignored for 24-hour time steps.

**References**

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

---

cal\_ET0\_from\_PM\_for\_daily

*Calculating reference evapotranspiration from Penman-Monteith for daily*

---

**Description**

Based on lat, z, J, Tmax, Tmin, n, RHmax, RHmin, windSpeed parameters, reference evapotranspiration was calculated by Penman-Monteith.

**Usage**

cal\_ET0\_from\_PM\_for\_daily(Latitude, Altitude, J, Tmax, Tmin, Rs, RHmean, Wind)

**Arguments**

Latitude	latitude (radian), positive for the northern hemisphere and negative for the southern hemisphere.
Altitude	station elevation above sea level [m].
J	is the number of the day in the year between 1 (1 January) and 365 or 366 (31 December).
Tmax	daily maximum air temperature (degrees Celsius).
Tmin	daily minimum air temperature (degrees Celsius).
Rs	Solar radiation [MJ m <sup>-2</sup> d <sup>-1</sup> ].
RHmean	daily mean relative humidity %.
Wind	wind speed at 2 m height [m s <sup>-1</sup> ].

**Value**

A vector for reference evapotranspiration (mm/day)

**References**

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

**Examples**

```
library(simET)
data("FIalfalfa")
names(FIalfalfa)
Result_data<- dplyr::mutate(FIalfalfa,
                           ET0=cal_ET0_from_PM_for_daily(Latitude=Latitude,
                                                           Altitude=Altitude,
                                                           J=Julian,
                                                           Tmax=Tmax,
                                                           Tmin=Tmin,
                                                           Rs=Rs,
                                                           RHmean=RHmean,
                                                           Wind=Wind))
names(Result_data)
```

---

cal\_ET0\_from\_PM\_for\_hourly

*Calculating reference evapotranspiration from Penman-Monteith method for hourly time step*

---

**Description**

Calculating reference evapotranspiration from Penman-Monteith method for hourly time step

**Usage**

```
cal_ET0_from_PM_for_hourly(
  slopVapourPressureCurve,
  netRadiation,
  soilHeatFlux,
  psychrometricConstant,
  meanHourlyTem,
  windSpeed,
  saturationVapourPressure,
  actualVapourPressure
)
```

**Arguments**

slopVapourPressureCurve      saturation slope vapour pressure curve at Thr [kPa &deg;C].

netRadiation      net radiation at the grass surface [MJ m-2 hour-1].

soilHeatFlux      soil heat flux density [MJ m-2 hour-1].

psychrometricConstant      psychrometric constant [kPa &deg;C].

meanHourlyTem      mean hourly air temperature [&deg;C].

windSpeed      average hourly wind speed [m s-1].

saturationVapourPressure      saturation vapour pressure at air temperature Thr [kPa].

actualVapourPressure      average hourly actual vapour pressure [kPa].

**Details**

In areas where substantial changes in wind speed, dewpoint or cloudiness occur during the day, calculation of the ETo equation using hourly time steps is generally better than using 24-hour calculation time steps. Such weather changes can cause 24-hour means to misrepresent evaporative power of the environment during parts of the day and may introduce error into the calculations. However, under most conditions, application of the FAO Penman-Monteith equation with 24-hour data produces accurate results.

**Value**

A vector for reference evapotranspiration [mm hour-1].

**Note**

With the advent of electronic, automated weather stations, weather data are increasingly reported for hourly or shorter periods. Therefore, in situations where calculations are computerized, the FAO Penman-Monteith equation can be applied on an hourly basis with good results. When applying the FAO Penman-Monteith equation on an hourly or shorter time scale, the equation and some of the procedures for calculating meteorological data should be adjusted for the smaller time step.

For the calculation of radiation parameters, see P74-75

**References**

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

---

cal\_extraterrestrialRadiation\_for\_daily

*Calculating extraterrestrial radiation for daily periods*

---

**Description**

The extraterrestrial radiation,  $R_a$ , for each day of the year and for different latitudes can be estimated from the solar constant, the solar declination and the time of the year.

**Usage**

cal\_extraterrestrialRadiation\_for\_daily(J, lat)

**Arguments**

J	is the number of the day in the year between 1 (1 January) and 365 or 366 (31 December).
lat	latitude (Radian), positive for the northern hemisphere and negative for the southern hemisphere.

**Value**

A vector for extraterrestrial radiation for daily(MJ m<sup>-2</sup> day<sup>-1</sup>)

**References**

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

---

cal\_extraterrestrialRadiation\_for\_shorter

*Calculating extraterrestrial radiation for hourly or shorter periods*

---

**Description**

Calculating extraterrestrial radiation for hourly or shorter periods

**Usage**

cal\_extraterrestrialRadiation\_for\_shorter(lat, J, t, lz, lm, t1)

**Arguments**

lat	latitude (radian), positive for the northern hemisphere and negative for the southern hemisphere.
J	is the number of the day in the year between 1 (1 January) and 365 or 366 (31 December).
t	standard clock time at the midpoint of the period (hour). For example for a period between 14.00 and 15.00 hours, t = 14.5.
lz	longitude of the centre of the local time zone (degrees west of Greenwich). For example, Lz = 75, 90, 105 and 120° for the Eastern, Central, Rocky Mountain and Pacific time zones (United States) and Lz = 0° for Greenwich, 330° for Cairo (Egypt), and 255° for Bangkok (Thailand), radian.
lm	longitude of the measurement site (degrees west of Greenwich) radian.
t1	length of the calculation period (hour)

**Value**

A vector for extraterrestrial Radiation (MJ m<sup>-2</sup> hour<sup>-1</sup>)

**References**

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

---

cal\_frictionVelocity *Calculating friction velocity*

---

**Description**

Calculating friction velocity

**Usage**

cal\_frictionVelocity(k = 0.4, zr = 2, u\_zr, d, z0)

**Arguments**

k	the von Karman constant (0.4)
zr	the height of the weather station(m).
u_zr	the wind speed(m/s) at the reference height zr (m).
d	zero plane displacement height (m)
z0	the crop roughness length(m)

**Value**

A vector for friction velocity(m/s)



---

cal_hourAngle	<i>Calculating hour angle</i>
---------------	-------------------------------

---

**Description**

Calculating hour angle

**Usage**

```
cal_hourAngle(th)
```

**Arguments**

th is the local solar time.

**Value**

A vector for hour angle (Radian)

**References**

Teh CBS. Introduction to mathematical modeling of crop growth: How the equations are derived and assembled into a computer model. Brown Walker Press, 2006.

**Examples**

```
cal_hourAngle(12)
```

---

cal_inverseRelativeDistance_Earth_sun	<i>Calculating inverse relative distance Earth-sun</i>
---------------------------------------	--

---

**Description**

Calculating inverse relative distance Earth-sun

**Usage**

```
cal_inverseRelativeDistance_Earth_sun(J)
```

**Arguments**

J is the number of the day in the year between 1 (1 January) and 365 or 366 (31 December).

**Value**

A vector for inverse relative distance Earth-sun (Radian)

**References**

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

cal\_Kcend\_for\_singleKc

*Crop coefficient for the end of the late season stage*

**Description**

Typical values for the crop coefficient at the end of the late season growth stage, Kc end, are listed in Table 12 for various agricultural crops.

**Usage**

cal\_Kcend\_for\_singleKc(RHmine, u2e, Ktable, he)

**Arguments**

RHmine	mean value for daily minimum relative humidity during the mid-season growth stage , for $20 \leq \text{RHmine} \leq 80$
u2e	mean value for daily wind speed at 2 m height over grass during the mid- season growth stage (m/s), for $1 \leq u2e \leq 6$
Ktable	value for Kc mid taken from Table 12
he	mean plant height during the mid-season stage (m) for $0.1 \text{ m} < h < 10 \text{ m}$

**Value**

A value for Kcend value

**Note**

only applied when the tabulated values for Kc end exceed 0.45. The equation reduces the Kc end with increasing RHmin. This reduction in Kc end is characteristic of crops that are harvested 'green' or before becoming completely dead and dry (Kc end  $\geq 0.45$ ).

No adjustment is made when Kc end (Table)  $< 0.45$  (Kc end = Kc end (Tab)). When crops are allowed to senesce and dry in the field (as evidenced by Kc end  $< 0.45$ ), u2 and RHmin have less effect on Kc end and no adjustment is necessary.

**References**

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

---

 cal\_Kcini\_for\_SingleKc

*Calculating Kcini value*


---

**Description**

Calculating Kcini value

**Usage**

cal\_Kcini\_for\_SingleKc(Pmean, ET0, tw, type, fw)

**Arguments**

Pmean	is the average depth of infiltrated water per wetting events(mm)
ET0	mean ET0 during initial period(mm/day)
tw	is the mean interval between wetting events(days)
type	soil type:"coarse soil textures" and "medium and fine soil textures"
fw	the fraction of surfaces wetted by irrigation or rain (0-1)

**Value**

A value for Kcini value

**References**

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

---

 cal\_Kcmid\_for\_singleKc

*Crop coefficient for the mid-season stage*


---

**Description**

Typical values for the crop coefficient at the end of the late season growth stage, Kc end, are listed in Table 12 for various agricultural crops. For specific adjustment in climates where RHmin differs from 45 % or where u2 is larger or smaller than 2.0 m/s.

**Usage**

cal\_Kcmid\_for\_singleKc(RHmine, u2e, Ktable, he)

**Arguments**

RHmine	mean value for daily minimum relative humidity during the mid-season growth stage , for $20 \leq \text{RHmine} \leq 80$
u2e	mean value for daily wind speed at 2 m height over grass during the mid- season growth stage (mls), for $1 \leq u2e \leq 6$
Ktable	value for Kc mid taken from Table 12
he	mean plant height during the mid-season stage [m] for $0.1 \text{ m} < h < 10 \text{ m}$

**Value**

A value for Kcmid value

**References**

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

---

cal\_Kc\_max\_for\_DualKc *An upper limit on the evaporation and transpiration from any cropped surface*

---

**Description**

It is imposed to reflect the natural constraints placed on available energy represented by the energy balance difference  $R_n - G - H$

**Usage**

cal\_Kc\_max\_for\_DualKc(u2, RHmin, h, Kcb)

**Arguments**

u2	The wind speed at 2 m
RHmin	Minimum relative humidity
h	Plant height
Kcb	Basal crop coefficient

**Value**

A vector for the upper limit on the evaporation and transpiration from any cropped surface

**References**

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

---

cal\_Kr\_for\_DualKc      *Dimensionless evaporation reduction coefficient*

---

### Description

It dependent on the soil water depletion (cumulative depth of evaporation) from the topsoil layer (Kr = 1 when De,i-1 is equal or lesser than REW)

### Usage

cal\_Kr\_for\_DualKc(TEW, REW, De)

### Arguments

TEW	maximum cumulative depth of evaporation (depletion) from the soil surface layer when Kr = 0 (TEW = total evaporable water)
REW	cumulative depth of evaporation (depletion) at the end of stage 1 (REW = readily evaporable water), mm
De	cumulative depth of evaporation (depletion) from the soil surface layer at the end of day i-1 (the previous day),mm

### Value

A value for evaporation reduction coefficient

### References

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

---

cal\_latentHeatFluxesForCrop  
*Calculating latent heat fluxes for crop*

---

### Description

Calculating latent heat fluxes for crop

### Usage

cal\_latentHeatFluxesForCrop(Delta, Ac, rho\_cp, D0, r\_c\_a, gamma, r\_c\_s)

**Arguments**

DELTA	the slope of the saturated vapor pressure curve (mbar K-1)
Ac	energy available to the crop (W m-2 ground)
rho_cp	the volumetric heat capacity for air (1221.09 J m-3 K-1)
D0	the vapor pressure deficit at the mean canopy flow
r_c_a	is the bulk boundary layer resistance (s m-1)
gamma	is the psychometric constant (0.658 mbar K-1)
r_c_s	the canopy resistance (s m-1)

**Value**

A vector for latent heat fluxes for crop (W m-2 ground)

---

cal\_latentHeatFluxesForSoil

*Calculating latent heat fluxes for soil*

---

**Description**

Calculating latent heat fluxes for soil

**Usage**

```
cal_latentHeatFluxesForSoil(
  DELTA,
  As,
  rho_cp = 1221.09,
  D0,
  r_s_a,
  gamma = 0.658,
  r_s_s
)
```

**Arguments**

DELTA	the slope of the saturated vapor pressure curve (mbar K-1)
As	energy available to the soil (W m-2 ground)
rho_cp	the volumetric heat capacity for air (1221.09 J m-3 K-1)
D0	the vapor pressure deficit at the mean canopy flow
r_s_a	the aerodynamic resistance between the soil and mean canopy flow (s m-1)
gamma	the psychometric constant (0.658 mbar K-1)
r_s_s	soil surface resistance, (s m-1)

**Value**

A vector for soil latent heat fluxes (W m-2 ground)

---

cal\_localDolarTime     *Calculating local solar time*

---

**Description**

Local solar time is different with local time.

**Usage**

```
cal_localDolarTime(td, t, gamma, gamma_sm)
```

**Arguments**

td	The day of year.
t	is the local time.
gamma	is the local longitude (Radian).
gamma_sm	is the standard longitude (Radian).

**Value**

A vector for local solar time(Hour)

**References**

Teh CBS.Introduction to mathematical modeling of crop growth: How the equations are derived and assembled into a computer model. Brown Walker Press, 2006.

**Examples**

```
cal_localDolarTime(td=1, t=12, gamma=0.52, gamma_sm=2.09)
```

---

cal\_meanCanopyFlowToReferenceLevel  
*Calculating mean canopy flow to reference level*

---

**Description**

Calculating mean canopy flow to reference level

**Usage**

```
cal_meanCanopyFlowToReferenceLevel(k = 0.4, u_, zr, d, h, nK, z0)
```

**Arguments**

k	the von Karman constant (0.4)
u_	the friction velocity (m s-1)
zr	is the reference height (m).the height of the weather station(m).
d	zero plane displacement height (m)
h	the plant height(m)
nK	the eddy diffusivity extinction coefficient (taken as 2)
z0	the crop roughness length (m)

**Value**

A vector for mean canopy flow to reference level

---

cal\_meanSaturationVapourPressure

*Calculating mean saturation vapour pressure*

---

**Description**

Due to the non-linearity of the above equation, the mean saturation vapour pressure for a day, week, decade or month should be computed as the mean between the saturation vapour pressure at the mean daily maximum and minimum air temperatures for that period.

**Usage**

cal\_meanSaturationVapourPressure(Tmax, Tmin)

**Arguments**

Tmax	the daily maximum air temperature(degrees Celsius).
Tmin	the daily minimum air temperature(degrees Celsius).

**Details**

Using mean air temperature instead of daily minimum and maximum temperatures results in lower estimates for the mean saturation vapour pressure. The corresponding vapour pressure deficit (a parameter expressing the evaporating power of the atmosphere) will also be smaller and the result will be some underestimation of the reference crop evapotranspiration. Therefore, the mean saturation vapour pressure should be calculated as the mean between the saturation vapour pressure at both the daily maximum and minimum air temperature.

**Value**

A vector for mean saturation vapour pressure (es)



## References

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

---

cal\_netLongwaveRadiation

*Calculating net longwave radiation Rnl*

---

## Description

Calculating net longwave radiation Rnl

## Usage

cal\_netLongwaveRadiation(TKmax, TKmin, ea, Rs, Rso)

## Arguments

TKmax	maximum absolute temperature during the 24-hour period [K].
TKmin	minimum absolute temperature during the 24-hour period [K].
ea	actual vapour pressure [kPa].
Rs	measured or calculated solar radiation [MJ m <sup>-2</sup> day <sup>-1</sup> ]. From cal_solarRadiation().
Rso	calculated clear-sky radiation [MJ m <sup>-2</sup> day <sup>-1</sup> ]. From cal_skySolarRadiation_withas_bs() or cal_skySolarRadiation_withas_elevation().

## Value

A vector for net outgoing longwave radiation [MJ m<sup>-2</sup> day<sup>-1</sup>]

## Note

The rate of longwave energy emission is proportional to the absolute temperature of the surface raised to the fourth power. This relation is expressed quantitatively by the Stefan-Boltzmann law. The net energy flux leaving the earth's surface is, however, less than that emitted and given by the Stefan-Boltzmann law due to the absorption and downward radiation from the sky. Water vapour, clouds, carbon dioxide and dust are absorbers and emitters of longwave radiation. Their concentrations should be known when assessing the net outgoing flux. As humidity and cloudiness play an important role, the Stefan-Boltzmann law is corrected by these two factors when estimating the net outgoing flux of longwave radiation. It is thereby assumed that the concentrations of the other absorbers are constant. An average of the maximum air temperature to the fourth power and the minimum air temperature to the fourth power is commonly used in the Stefan-Boltzmann equation for 24-hour time steps. The term  $(0.34-0.14*\sqrt{ea})$  expresses the correction for air humidity, and will be smaller if the humidity increases. The effect of cloudiness is expressed by  $(1.35 Rs/Rso - 0.35)$ . The term becomes smaller if the cloudiness increases and hence Rs decreases. The smaller the correction terms, the smaller the net outgoing flux of longwave radiation. Note that the Rs/Rso term in Equation 39 must be limited so that  $Rs/Rso \leq 1.0$ . Where measurements of incoming and outgoing short and longwave radiation during bright sunny and overcast hours are available, calibration of the coefficients in Equation 39 can be carried out.

**References**

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

---

cal\_netRadiation      *Calculating net radiation Rn*

---

**Description**

The net radiation (Rn) is the difference between the incoming net shortwave radiation (Rns) and the outgoing net longwave radiation (Rnl).

**Usage**

cal\_netRadiation(Rns, Rnl)

**Arguments**

Rns                    incoming net shortwave radiation. From cal\_netSolarRadiation().  
 Rnl                    outgoing net longwave radiation. From cal\_netLongwaveRadiation().

**Value**

A vector for net radiation

**References**

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

---

cal\_netRadiationForCrop  
                                  *Calculating net radiation available to the crop*

---

**Description**

Calculating net radiation available to the crop

**Usage**

cal\_netRadiationForCrop(pRn, Rn)

**Arguments**

pRn                    canopy penetration probability for net radiation  
 Rn                    the net radiation (W/m<sup>2</sup> ground).see cal\_hourlyNetRadiation()

**Value**

A vector for net radiation available to the crop (W/m<sup>2</sup> ground)

---

cal\_netRadiationForSoil

*Calculating net radiation available to the soil*

---

**Description**

Calculating net radiation available to the soil

**Usage**

cal\_netRadiationForSoil(pRn, Rn, G)

**Arguments**

pRn	canopy penetration probability for net radiation
Rn	the net radiation (W/m <sup>2</sup> ground).see cal_hourlyNetRadiation()
G	the soil heat flux(W/m <sup>2</sup> ground)

**Value**

A vector for net radiation available to the soil (W/m<sup>2</sup> ground)

---

cal\_netRadiationForSystem

*Calculating net radiation available to the system(soil and crop)*

---

**Description**

Calculating net radiation available to the system(soil and crop)

**Usage**

cal\_netRadiationForSystem(Ac, As)

**Arguments**

Ac	net radiation available to the crop (W/m <sup>2</sup> ground)
As	net radiation available to the soil(W/m <sup>2</sup> ground)

**Value**

A vector for net radiation available to the system(soil and crop) (W/m<sup>2</sup> ground)

---

cal\_netSolarRadiation *Calculating net solar (shortwave radiation) Rns*

---

### Description

The net shortwave radiation resulting from the balance between incoming and reflected solar radiation.

### Usage

cal\_netSolarRadiation(alpha, Rs)

### Arguments

alpha	albedo or canopy reflection coefficient, which is 0.23 for the hypothetical grass reference crop [dimensionless].
Rs	the incoming solar radiation [MJ m <sup>-2</sup> day <sup>-1</sup> ]. From cal_solarRadiation()

### Value

A vector for net solar or shortwave radiation [MJ m<sup>-2</sup> day<sup>-1</sup>].

### References

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

---

cal\_percolationForExcessWater  
*Calculating percolation for excess water*

---

### Description

Calculating percolation for excess water

### Usage

cal\_percolationForExcessWater(THETA\_i\_t0, Pe\_i\_t1, THETA\_sat\_i)

### Arguments

THETA_i_t0	the water amount of the day before in soil layer i (mm)
Pe_i_t1	the percolation of previous soil layer(mm)
THETA_sat_i	soil saturation water amount (mm)

**Value**

A value for percolation for excess water (mm)

---

cal\_psychrometricConstant

*Calculating psychrometric constant*

---

**Description**

Calculating psychrometric constant

**Usage**

cal\_psychrometricConstant(atmospheric\_pressure)

**Arguments**

atmospheric\_pressure

atmospheric pressure (kPa).

**Value**

A vector for Psychrometric constant (kPa/degree Celsius)

**References**

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

**Examples**

cal\_psychrometricConstant(100.1235)

---

cal\_reductionFactorForE

*Calculating reduction factor for evaporation*

---

**Description**

Calculating reduction factor for evaporation

**Usage**

cal\_reductionFactorForE(THETA, THETA\_sat)

**Arguments**

THETA	the water amount of the day before in stop soil layer (mm)
THETA_sat	soil saturation water content in stop oil layer (mm)

**Value**

A value for reduction factor for evaporation

---

cal\_reductionFactorForT  
*calculating reduction factor for transpiration*

---

**Description**

calculating reduction factor for transpiration

**Usage**

cal\_reductionFactorForT(THETA\_v\_wp, p, THETA\_v\_sat, THETA\_v)

**Arguments**

THETA_v_wp	soil water content at wilting point (m <sup>3</sup> m <sup>-3</sup> )
p	a coefficient. 0.5 for C3 and 0.3 for C4 plant
THETA_v_sat	soil saturation water content (m <sup>3</sup> m <sup>-3</sup> )
THETA_v	the water amount of the day before in root layer(m <sup>3</sup> m <sup>-3</sup> )

**Value**

A value for reduction factor for transpiration

---

cal\_relativeHumidity *Calculating relative humidity*

---

**Description**

The relative humidity (RH) expresses the degree of saturation of the air as a ratio of the actual (ea) to the saturation (eo(T)) vapour pressure at the same temperature (T).

**Usage**

cal\_relativeHumidity(ea, e0)

**Arguments**

ea	actual saturation vapour pressure. From cal_ActualVapourPressure_for_*
e0	saturation vapour pressure. From cal_saturationVapourPressure()

**Details**

Relative humidity is the ratio between the amount of water the ambient air actually holds and the amount it could hold at the same temperature. It is dimensionless and is commonly given as a percentage. Although the actual vapour pressure might be relatively constant throughout the day, the relative humidity fluctuates between a maximum near sunrise and a minimum around early afternoon (Figure 12). The variation of the relative humidity is the result of the fact that the saturation vapour pressure is determined by the air temperature. As the temperature changes during the day, the relative humidity also changes substantially.

**Value**

A vector for relative humidity %

**References**

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

---

cal_Rs_from_Na	<i>Calculating Solar radiation from actual duration of sunshine</i>
----------------	---

---

**Description**

Calculating Solar radiation from actual duration of sunshine

**Usage**

```
cal_Rs_from_Na(as = 0.25, bs = 0.5, Na, Latitude, J)
```

**Arguments**

as	regression constant, expressing the fraction of extraterrestrial radiation reaching the earth on overcast days (n = 0). Default is 0.25.
bs	as+bs is fraction of extraterrestrial radiation reaching the earth on clear days (n = N). Default is 0.50.
Na	actual duration of sunshine [hour].
Latitude	latitude (angert).
J	is the number of the day in the year between 1 (1 January) and 365 or 366 (31 December).

**Value**

A vector for solar radiation(MJ m<sup>-2</sup> d<sup>-1</sup>)

**References**

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

---

cal\_saturationVapourPressure

*Calculating saturation vapour pressure*

---

**Description**

saturation vapour pressure at the air temperature T.

**Usage**

cal\_saturationVapourPressure(Tem)

**Arguments**

Tem                      air temperature (degrees Celsius).

**Value**

A vector for saturation vapour pressure at the air temperature T (kPa).

**References**

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

---

cal\_sensibleHeatFluxesForCrop

*Calculating sensible heat fluxes for crop*

---

**Description**

Calculating sensible heat fluxes for crop

**Usage**

cal\_sensibleHeatFluxesForCrop(gamma, Ac, r\_c\_s, r\_c\_a, rho\_cp, D0, DELTA)



**Arguments**

gamma	is the psychometric constant (0.658 mbar K-1)
Ac	energy available to the crop (W m-2 ground)
r_c_s	the canopy resistance (s m-1)
r_c_a	is the bulk boundary layer resistance (s m-1)
rho_cp	the volumetric heat capacity for air (1221.09 J m-3 K-1)
D0	the vapor pressure deficit at the mean canopy flow
DELTA	the slope of the saturated vapor pressure curve (mbar K-1)

**Value**

A vector for sensible heat fluxes for crop

---

cal\_sensibleHeatFluxesForSoil  
*Calaulating sensible heat fluxes for soil*

---

**Description**

Calaulating sensible heat fluxes for soil

**Usage**

```
cal_sensibleHeatFluxesForSoil(
  gamma = 0.659,
  As,
  r_s_s,
  r_s_a,
  rho_cp,
  D0,
  DELTA
)
```

**Arguments**

gamma	is the psychometric constant (0.658 mbar K-1)
As	energy available to the soil (W m-2 ground)
r_s_s	soil surface resistance, (s m-1)
r_s_a	is the aerodynamic resistance between the soil and mean anopy flow (s m-1);
rho_cp	the volumetric heat capacity for air (1221.09 J m-3 K-1)
D0	the vapor pressure deficit at the mean canopy flow
DELTA	the slope of the saturated vapor pressure curve (mbar K-1)

**Value**

A vector for soil sensible heat fluxes(W m-2 ground)

---

cal\_skySolarRadiation\_withas\_bs

*Calculating clear sky solar radiation with as and bs*

---

### Description

The calculation of the clear-sky radiation,  $R_{so}$ , when  $n = N$ , is required for computing net longwave radiation.

### Usage

cal\_skySolarRadiation\_withas\_bs(as, bs, Ra)

### Arguments

as, bs	as+bs fraction of extraterrestrial radiation reaching the earth on clear-sky days ( $n = N$ ).
Ra	extraterrestrial radiation [ $\text{MJ m}^{-2} \text{ day}^{-1}$ ]. From cal_extraterrestrialRadiation_for_daily()

### Value

A vector for clear-sky solar radiation [ $\text{MJ m}^{-2} \text{ day}^{-1}$ ].

### References

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

---

cal\_skySolarRadiation\_withas\_elevation

*Calculating clear sky solar radiation with elevation*

---

### Description

The calculation of the clear-sky radiation,  $R_{so}$ , when  $n = N$ , is required for computing net longwave radiation.

### Usage

cal\_skySolarRadiation\_withas\_elevation(z, Ra)

### Arguments

z	station elevation above sea level [m].
Ra	extraterrestrial radiation [ $\text{MJ m}^{-2} \text{ day}^{-1}$ ]. From cal_extraterrestrialRadiation_for_daily()

**Value**

A vector for clear-sky solar radiation [MJ m<sup>-2</sup> day<sup>-1</sup>].

**References**

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

---

cal\_slopeOfSaturationVapourPressureCurve

*Calculating slope of saturation vapour pressure curve*

---

**Description**

Calculating slope of saturation vapour pressure curve

**Usage**

cal\_slopeOfSaturationVapourPressureCurve(Tem)

**Arguments**

Tem                    is air temperature (degrees Celsius).

**Details**

In the FAO Penman-Monteith equation, where it occurs in the numerator and denominator, the slope of the vapour pressure curve is calculated using mean air temperature.

**Value**

A vector for slope of saturation vapour pressure curve at air temperature T

**References**

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

---

cal\_soilHeatFlux      *Calculating Soil/ground heat flux*

---

**Description**

Calculating Soil/ground heat flux

**Usage**

cal\_soilHeatFlux(pRn, Rn)

**Arguments**

pRn                    the canopy penetration probability for net radiation  
Rn                     the net radiation (W/m<sup>2</sup> ground)

**Value**

A vector for the soil heat flux (W/m<sup>2</sup> ground)

---

cal\_soilHeatFlux\_day      *Calculating soil heat flux(G) for day/ten-day periods*

---

**Description**

As the magnitude of the day or ten-day soil heat flux beneath the grass reference surface is relatively small, it may be ignored .

**Usage**

cal\_soilHeatFlux\_day()

**Value**

A value for 0

**References**

FAO Irrigation and drainage paper 56 (P54)

---

`cal_soilHeatFlux_general`*Calculating soil heat flux (G) for general*

---

**Description**

Complex models are available to describe soil heat flux. Because soil heat flux is small compared to  $R_n$ , particularly when the surface is covered by vegetation and calculation time steps are 24 hours or longer, a simple calculation procedure is presented here for long time steps, based on the idea that the soil temperature follows air temperature.

**Usage**`cal_soilHeatFlux_general(cs, T1, T0, delta_t, delta_z)`**Arguments**

<code>cs</code>	soil heat capacity [MJ m <sup>-3</sup> degrees Celsius <sup>-1</sup> ].
<code>T1</code>	air temperature at time $i$ [degrees Celsius].
<code>T0</code>	air temperature at time $i-1$ [degrees Celsius].
<code>delta_t</code>	length of time interval [day].
<code>delta_z</code>	effective soil depth [m].

**Value**

A vector for soil heat flux [MJ m<sup>-2</sup> day<sup>-1</sup>]

**Note**

Complex models are available to describe soil heat flux. Because soil heat flux is small compared to  $R_n$ , particularly when the surface is covered by vegetation and calculation time steps are 24 hours or longer, a simple calculation procedure is presented here for long time steps, based on the idea that the soil temperature follows air temperature.

**References**

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

---

 cal\_soilHeatFlux\_hourly

*Calculating soil heat flux(G) for hourly/shorter periods*


---

**Description**

For hourly (or shorter) calculations, G beneath a dense cover of grass does not correlate well with air temperature.

**Usage**

```
cal_soilHeatFlux_hourly(Rn, periods)
```

**Arguments**

Rn	net radiation.From cal_netRadiation().
periods	"daylight" or "nighttime".

**Value**

A vector for soil heat flux [MJ m<sup>-2</sup> day<sup>-1</sup>]

**Note**

Where the soil is warming, the soil heat flux G is positive. The amount of energy required for this process is subtracted from Rn when estimating evapotranspiration.

**References**

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

---

 cal\_soilHeatFlux\_monthly

*Calculating soil heat flux(G) for monthly periods*


---

**Description**

When assuming a constant soil heat capacity of 2.1 MJ m<sup>-3</sup> °C<sup>-1</sup> and an appropriate soil depth, cal\_soilHeatFlux\_general can be used to derive G for monthly periods.

**Usage**

```
cal_soilHeatFlux_monthly(T1, T0, Tmonth2 = TRUE)
```

**Arguments**

T1	air temperature at time i [degrees Celsius].
T0	air temperature at time i-1 [degrees Celsius].
Tmonth2	Is the mean air temperature of next month know?

**Value**

A vector for soil heat flux [MJ m<sup>-2</sup> day<sup>-1</sup>]

**References**

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

---

cal\_soilSurfaceResistance

*Calculating soil surface resistance*

---

**Description**

Calculating soil surface resistance

**Usage**

```
cal_soilSurfaceResistance(
  tau,
  l,
  PHI_p,
  Dm_v,
  lambda_p,
  THETA_v_l,
  THETA_v_sat_l
)
```

**Arguments**

tau	soil tortuosity (taken as 2)
l	is the dry soil layer thickness (taken as the first soil layer thickness) (m)
PHI_p	is soil porosity
Dm_v	the vapor diffusion coefficient in air (24.7 10 <sup>-6</sup> m <sup>2</sup> s <sup>-1</sup> )
lambda_p	the soil pore-size distribution index from the Brooks-Corey equation.
THETA_v_l	volumetric soil water content (m <sup>3</sup> m <sup>-3</sup> ) of the first soil layer
THETA_v_sat_l	saturated soil water content (m <sup>3</sup> m <sup>-3</sup> ) of the first soil layer

**Value**

A vector for the soil surface resistance (s m<sup>-1</sup>)

---

cal\_soilSurfaceToMeanCanopyFlow

*Calculating soil surface to mean canopy flow*

---

### Description

the resistance between the soil surface and the mean canopy flow (s m-1)

### Usage

cal\_soilSurfaceToMeanCanopyFlow(h, nK, Kh, zs0, z0, d)

### Arguments

h	the plant height(m)
nK	the eddy diffusivity extinction coefficient (taken as 2)
Kh	eddy diffusivity at the canopy top(m <sup>2</sup> /s)
zs0	is the soil surface roughnesslength (m). Note: for flat, tilled land, zs0 can be taken as 0.004 m.
z0	the crop roughness length (m)
d	zero plane displacement height (m)

### Value

A vector for aerodynamic resistancesoil surface to mean canopy flow (s m-1)

---

cal\_solarDeclination *Calculating solar declination*

---

### Description

Calculating solar declination

### Usage

cal\_solarDeclination(td)

### Arguments

td	is the day of year.
----	---------------------

### Value

A vector for solar declination (Radian)



### Note

The solar declination actually varies throughout the day too but its variation is very small; thus, it is often ignored. Negative angles occur when the angle is below the equator plane, positive for above the equator.

### References

Teh CBS. Introduction to mathematical modeling of crop growth: How the equations are derived and assembled into a computer model. Brown Walker Press, 2006.

### Examples

```
cal_solarDeclination(34)
```

---

```
cal_solarDeclination_in_FAO
```

*Calculating solar declination with FAO56 method*

---

### Description

Calculating solar declination with FAO56 method

### Usage

```
cal_solarDeclination_in_FAO(J)
```

### Arguments

J is the number of the day in the year between 1 (1 January) and 365 or 366 (31 December)

### Value

A vector for solar declination(Radian)

### References

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

---

cal\_solarInclination    *Calculating solar inclination*

---

**Description**

A parameter used to determine the position of the sun relative to the observer (the other one is the angle from south). Conversion relationship with solar altitude angle: solar inclination= $\pi/2$ -solar altitude.

**Usage**

```
cal_solarInclination(solar_declination, latitude, hour_anger)
```

**Arguments**

solar\_declination    is solar declination anger. It can be calculated from cal\_solardeclination().  
latitude    is the latitude data (Radian).  
hour\_anger    is hour anger. It can be calculated from cal\_hourangle().

**Value**

A vector for solar inclination (Radian)

**References**

Teh CBS. Introduction to mathematical modeling of crop growth: How the equations are derived and assembled into a computer model. Brown Walker Press, 2006.

**Examples**

```
cal_solarInclination(solar_declination=-0.297, latitude=30, hour_anger=0)
```

---

cal\_solarRadiation    *Calculating Solar radiation*

---

**Description**

If the solar radiation,  $R_s$ , is not measured, it can be calculated with the Angstrom formula, which relates solar radiation to extraterrestrial radiation and relative sunshine duration. This is a shortwave radiation.

**Usage**

```
cal_solarRadiation(as = 0.25, bs = 0.5, n, N, Ra)
```

**Arguments**

as	regression constant, expressing the fraction of extraterrestrial radiation reaching the earth on overcast days ( $n = 0$ ). Default is 0.25.
bs	as+bs is fraction of extraterrestrial radiation reaching the earth on clear days ( $n = N$ ). Default is 0.50.
n	actual duration of sunshine [hour].
N	maximum possible duration of sunshine or daylight hours [hour]. from cal_daylightHours()
Ra	extraterrestrial radiation [ $\text{MJ m}^{-2} \text{ day}^{-1}$ ]. From cal_extraterrestrialRadiation_for_daily()

**Value**

A vector for solar or shortwave radiation [ $\text{MJ m}^{-2} \text{ day}^{-1}$ ]

**Note**

$R_s$  is expressed in the above equation in  $\text{MJ m}^{-2} \text{ day}^{-1}$ . The corresponding equivalent evaporation in  $\text{mm day}^{-1}$  is obtained by multiplying  $R_s$  by 0.408 (Equation 20). Depending on atmospheric conditions (humidity, dust) and solar declination (latitude and month), the Angstrom values as and bs will vary. Where no actual solar radiation data are available and no calibration has been carried out for improved as and bs parameters, the values as = 0.25 and bs = 0.50 are recommended.

**References**

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

---

cal\_sunsetHourAngle     *Calculating sunset hour angle*

---

**Description**

Calculating sunset hour angle

**Usage**

cal\_sunsetHourAngle(lat, solar\_declination)

**Arguments**

lat	latitude (Radian), positive for the northern hemisphere and negative for the southern hemisphere.
solar_declination	solar declination(Radina).

**Value**

A vector for sunset Hour Angle(Radian)

**References**

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

---

cal_sunsetTime	<i>Calculating the local solar time for sunset/sunrise</i>
----------------	--

---

**Description**

Calculating the local solar time for sunset/sunrise.

**Usage**

```
cal_sunsetTime(solar_declination, latitude)
```

**Arguments**

solar\_declination  
can be calculated by cal\_solardeclination().

latitude  
is latitude data(Radian).

**Details**

Knowing the time of sunrise can calculate the time of sunset.  $\text{sunrise\_time}=24-\text{sunset\_time}$ .  $\text{Day\_length}=2*(\text{sunset\_time}-12)$ .

**Value**

A vector for the local solar time for sunset/sunrise

**References**

Teh CBS. Introduction to mathematical modeling of crop growth: How the equations are derived and assembled into a computer model. Brown Walker Press, 2006.

---

cal\_TemMean                    *calculating the mean daily air temperature*

---

**Description**

calculating the mean daily air temperature

**Usage**

cal\_TemMean(Tmax, Tmin)

**Arguments**

Tmax                    the daily maximum.The temperature is given in degree Celsius or Fahrenheit.  
Tmin                    the daily minimum.The temperature is given in degree Celsius,or Fahrenheit.

**Details**

It is only employed in the FAO Penman-Monteith equation to calculate the slope of the saturation vapor pressure curves and the impact of mean air density as the effect of temperature variations on the value of the climatic parameter is small in these cases. For standardization, Tmean for 24-hour periods is defined as the mean of mean of the daily maximum and minimum temperatures rather than as the average of hourly temperature measurements.

**Value**

A vector for the mean daily air temperature

**References**

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

---

cal\_TEW\_for\_DualKc            *calculating total evaporable water*

---

**Description**

maximum depth of water that can be evaporated from the soil when the topsoil has been initially completely wetted. Estimated TEW for Kr calculation

**Usage**

cal\_TEW\_for\_DualKc(FC, WP, Ze)

**Arguments**

FC	Soil water content at field capacity, m <sup>3</sup> m <sup>-3</sup>
WP	Soil water content at wilting point, m <sup>3</sup> m <sup>-3</sup>
Ze	Depth of the surface soil layer that is subject to drying by way of evaporation, 0.10-0.15 m.

**Value**

A value for total evaporable water

**References**

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

---

cal\_totalLatentHeatFlux

*Calculating total latent heat flux*

---

**Description**

Calculating total latent heat flux

**Usage**

```
cal_totalLatentHeatFlux(
  DELTA,
  gamma,
  r_a_a,
  r_c_a,
  r_s_a,
  r_c_s,
  r_s_s,
  A,
  rho_cp = 1221.09,
  D,
  As,
  Ac
)
```

**Arguments**

DELTA	the slope of the saturated vapor pressure curve (mbar K <sup>-1</sup> )
gamma	is the psychometric constant (0.658 mbar K <sup>-1</sup> )
r_a_a	the aerodynamic resistance between the mean canopy flow and reference height (s m <sup>-1</sup> )

r_c_a	the bulk boundary layer resistance (s m-1)
r_s_a	is the aerodynamic resistance between the soil and mean canopy flow (s m-1)
r_c_s	the canopy resistance(s m-1)
r_s_s	soil surface resistance (s m-1)
A	energy available to the system (total)(W m-2 ground)
rho_cp	is the volumetric heat capacity for air (1221.09 J m-3 K-1)
D	the vapor pressure deficit (mbar)
As	energy available to soil (W m-2 ground)
Ac	energy available to crop (W m-2 ground)

**Value**

A vector for the total latent heat flux (W m-2 ground)

---

cal\_WaterStressCoef    *Calculating water stress coefficient*

---

**Description**

Calculating water stress coefficient

**Usage**

cal\_WaterStressCoef(Dr, TAW, p)

**Arguments**

Dr	root zone depletion(mm).
TAW	total available soil water in the root zone(mm).
p	fraction of TAW that a crop can extract from the root zone without suffering water stress.

**Value**

A value for water stress coefficient which is a dimensionless transpiration reduction factor dependent on available soil water

---

cal\_windSpeed\_Canopy *Calculating wind speed above and within the canopies.*

---

### Description

Calculating wind speed above and within the canopies.

### Usage

```
cal_windSpeed_Canopy(z, h, u_, k = 0.4, d, z0, nu = 2)
```

### Arguments

z	the height (m)
h	the plant height (m)
u_	the friction velocity (m/s)
k	the von Karman constant (0.4)
d	zero plane displacement height (m)
z0	the crop roughness length(m)
nu	the wind speed extinction coefficient (taken as 2)

### Value

A vector for the wind speed (m/s) at height z(m)

---

cal\_zeroPlaneHeight *Calculating zero plane displacement height*

---

### Description

Calculating zero plane displacement height

### Usage

```
cal_zeroPlaneHeight(h)
```

### Arguments

h	the plant height (m)
---	----------------------

### Value

A vector for zero plane displacement height (m)



---

compare\_model\_plot     *Show the results of different models*

---

**Description**

Show the results of different models

**Usage**

```
compare_model_plot(model_list, names)
```

**Arguments**

model\_list     List. Including output results of different models.  
names           Vector. Name of models.

**Value**

A list for ggplot2 plot

---

convert\_angert\_to\_radian  
*Converting angert to radian*

---

**Description**

Converting the unit of angle in longitude and latitude into the unit of radian.

**Usage**

```
convert_angert_to_radian(anger)
```

**Arguments**

anger           Longitude or dimension in Angle.

**Value**

A vector for longitude or dimension in radian.

**Examples**

```
convert_angert_to_radian(98.8)
```

---

convert\_Date\_to\_dayofyear  
*Convert date to day of year*

---

**Description**

Convert date to day of year

**Usage**

convert\_Date\_to\_dayofyear(Date)

**Arguments**

Date is a date format data.

**Value**

A vector for the day of year.

---

convert\_degreesCelsius\_to\_Fahrenheit  
*Convert degrees Celsius to Fahrenheit*

---

**Description**

Convert degrees Celsius to Fahrenheit

**Usage**

convert\_degreesCelsius\_to\_Fahrenheit(degrees\_Celsius)

**Arguments**

degrees\_Celsius  
temperature in degrees Celsius(°C).

**Value**

A vector for a temperature in Fahrenheit(°F).

**References**

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

---

convert\_Fahrenheit\_to\_degreesCelsius  
*Convert Fahrenheit to degrees Celsius*

---

**Description**

Convert Fahrenheit to degrees Celsius

**Usage**

convert\_Fahrenheit\_to\_degreesCelsius(Fahrenheit)

**Arguments**

Fahrenheit      temperature in Fahrenheit(°F).

**Value**

A vector for temperature in degrees Celsius(°C)

**References**

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

---

convert\_Rad\_unit      *Convert radiation unit*

---

**Description**

Type has the following types: MJ\_m2\_day\_to\_J\_cm2\_day; MJ\_m2\_day\_to\_cal\_cm2\_day; MJ\_m2\_day\_to\_W\_m2; cal\_cm2\_day\_to\_MJ\_m2\_day cal\_cm2\_day\_to\_J\_cm2\_day cal\_cm2\_day\_to\_W\_m2 cal\_cm2\_day\_to\_mm\_day W\_m2\_to\_MJ\_m2\_day W\_m2\_to\_J\_cm2\_day W\_m2\_to\_cal\_cm2\_day W\_m2\_to\_mm\_day mm\_day\_to\_MJ\_m2\_day mm\_day\_to\_J\_cm2\_day mm\_day\_to\_cal\_cm2\_day mm\_day\_to\_W\_m2

**Usage**

convert\_Rad\_unit(rad, type)

**Arguments**

rad                  Radiation data need to be converted from one unit to another unit.  
 type                Used to specify how to convert.

**Value**

A vector for radiation converted unit

## References

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

---

convert\_windSpeed\_to\_2m

*Convert wind speed to the standard of 2m*

---

## Description

For the calculation of evapotranspiration, wind. Speed measured at 2 m above the surface is required. To adjust wind speed data obtained from instruments placed at elevations other than the standard height of 2 m, a logarithmic wind speed profile may be used for measurements above a short grassed surface.

## Usage

convert\_windSpeed\_to\_2m(uz, z)

## Arguments

uz	measured wind speed at z m above ground surface [m s-1].
z	height of measurement above ground surface [m].

## Value

A vector for wind speed at 2 m above ground surface [m s-1].

## References

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

---

create\_modelDF

*Create a csv file or a dataframe in R to store the model data*

---

## Description

#' @title Converting the every sheets in XLSX file to csv files #' @param xlsx\_file the xlsx file path  
 #' @export #' @importFrom utils write.csv #' @return No return value

## Usage

create\_modelDF(TreNum = 1, rowNum = 1)

**Arguments**

TreNum            Number. Need to generate how many treatment column  
 rowNum            Number. The number of row.

**Details**

```
convert_xlsx_to_csv<-function(xlsx_file) # library(readxl) # library(stringr) #sheetsname Sheetnames<-
readxl::excel_sheets(xlsx_file)
```

```
#xlsx2csv
```

```
for (i in Sheetnames) data<-readxl::read_xlsx(xlsx_file,sheet=i) write.csv(data,file = stringr::str_c(i,".csv"),row.names
= FALSE)
```

Latitude and Longitude use radians as units; Altitude use 'm' as units; Na use 'hour' as units; Tmax and Tmin use Celsius as units; Wind use m/s as units; RHmean and RHmin use percent sign as units; Rs use MJ M-2 day-1 as units Height use cm as units; SoilWater and Irrigation use mm as units; GroundwaterDepth use cm as unit.

**Value**

A dataframe and a csv file (if to\_CSV\_file==TRUE)

**Note**

The column of soilwater refers to the measured soil water (mm) in the maximum root layer, which is used to compare the difference between the simulated value and the measured value, which is an optional variable. The columns of Stage includes four stages: Ini, Development, Mid, End, which are mainly determined by LAI and growth status and can refer to Allen et al., (1998).

---

estimate\_ea

*Estimating missing humidity data*

---

**Description**

Where humidity data are lacking or are of questionable quality, an estimate of actual vapour pressure, ea, can be obtained by assuming that dewpoint temperature (Tdew) is near the daily minimum temperature (Tmin). This statement implicitly assumes that at sunrise, when the air temperature is close to Tmin, that the air is nearly saturated with water vapour and the relative humidity is nearly 100 %.

**Usage**

```
estimate_ea(Tmin)
```

**Arguments**

Tmin            the minimum tem daily.

**Value**

A vector for humidity

**Note**

The relationship Tdew near Tmin holds for locations where the cover crop of the station is well watered. However, particularly for arid regions, the air might not be saturated when its temperature is at its minimum. Hence, Tmin might be greater than Tdew and a further calibration may be required to estimate dewpoint temperatures. In these situations, "Tmin" in the above equation may be better approximated by subtracting 2-3 degrees Celsius from Tmin. Appropriate correction procedures are given in Annex 6. In humid and subhumid climates, Tmin and Tdew measured in early morning may be less than Tdew measured during the daytime because of condensation of dew during the night. After sunrise, evaporation of the dew will once again humidify the air and will increase the value measured for Tdew during the daytime. This phenomenon is demonstrated in Figure 5.4 of Annex 5. However, it is standard practice in 24-hour calculations of ETo to use Tdew measured or calculated during early morning. The estimate for ea from Tmin should be checked. When the prediction by Equation 48 is validated for a region, it can be used for daily estimates of ea.

**References**

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

---

estimate\_ET0\_with\_TmaxAndTmin

*Estimating ETO with Tmax and Tmin*

---

**Description**

When solar radiation data, relative humidity data and/or wind speed data are missing, they should be estimated using the procedures presented in this section. As an alternative, ETo can be estimated using the Hargreaves ETo equation.

**Usage**

estimate\_ET0\_with\_TmaxAndTmin(Tmean, Tmax, Tmin, Ra)

**Arguments**

Tmean	mean temperature.
Tmax	max temperature.
Tmin	min temperature.
Ra	extraterrestrial radiation [mm day-1].

**Value**

A vector for reference evapotranspiration (mm day-1).

**Note**

Units for both ETo and Ra in Equation 52 are mm day<sup>-1</sup>. Equation 52 should be verified in each new region by comparing with estimates by the FAO Penman-Monteith equation (Equation 6) at weather stations where solar radiation, air temperature, humidity, and wind speed are measured. If necessary, Equation 52 can be calibrated on a monthly or annual basis by determining empirical coefficients where  $E_{To} = a + b E_{To}^{Eq.52}$ , where the Eq. 52 subscript refers to ETo predicted using Equation 52. The coefficients a and b can be determined by regression analyses or by visual fitting. In general, estimating solar radiation, vapor pressure and wind speed as described in Equations 48 to 51 and Table 4 and then utilizing these estimates in Equation 6 (the FAO Penman-Monteith equation) will provide somewhat more accurate estimates as compared to estimating ETo directly using Equation 52. This is due to the ability of the estimation equations to incorporate general climatic characteristics such as high or low wind speed or high or low relative humidity into the ETo estimate made using Equation 6. Equation 52 has a tendency to underpredict under high wind conditions ( $u_2 > 3$  m/s) and to overpredict under conditions of high relative humidity.

---

 estimate\_goodnessOfFit

*Calculating the goodness-of-fit indicators between measured and simulated values*

---

**Description**

Calculating the goodness-of-fit indicators between measured and simulated values

**Usage**

estimate\_goodnessOfFit(Sim, Obs)

**Arguments**

Sim	The simulation value of model.
Obs	The observed value.

**Value**

A vector for the goodness-of-fit indicators

---

 estimate\_LAI\_for\_alfalfa

*Estimate LAI for alfalfa*


---

### Description

Estimate LAI for alfalfa

### Usage

estimate\_LAI\_for\_alfalfa(hc)

### Arguments

hc                    is the vegetation height in meter. (in meter)

### Value

A vector for leaf are index of alfalfa

### References

Zhao C , Feng Z , Chen G . Soil water balance simulation of alfalfa (*Medicago sativa* L.) in the semiarid Chinese Loess Plateau[J]. *Agricultural Water Management*, 2004, 69(2):0-114.

---

 estimate\_Rs\_for\_islandLocations

*Estimating solar radiation for island locations*


---

### Description

For island locations, where the land mass has a width perpendicular to the coastline of 20 km or less, the air masses influencing the atmospheric conditions are dominated by the adjacent water body in all directions. The temperature method is not appropriate for this situation. Where radiation data from another location on the island are not available, a first estimate of the monthly solar average can be obtained from the empirical relation.

### Usage

estimate\_Rs\_for\_islandLocations(Ra, b = 4)

### Arguments

Ra                    extraterrestrial radiation [MJ m<sup>-2</sup> day<sup>-1</sup>].  
 b                      empirical constant, equal to 4 MJ m<sup>-2</sup> day<sup>-1</sup>.



**Value**

A vector for solar radiation

**Note**

This relationship is only applicable for low altitudes (from 0 to 100 m). The empirical constant represents the fact that in island locations some clouds are usually present, thus making the mean solar radiation 4 MJ m<sup>-2</sup> day<sup>-1</sup> below the nearly clear sky envelope (0.7 Ra). Local adjustment of the empirical constant may improve the estimation. The method is only appropriate for monthly calculations. The constant relation between Rs and Ra does not yield accurate daily estimates.

**References**

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

---

estimate\_Rs\_from\_airTemDiff

*Estimating solar radiation data derived from air temperature differences*

---

**Description**

The difference between the maximum and minimum air temperature is related to the degree of cloud cover in a location. Clear-sky conditions result in high temperatures during the day (Tmax) because the atmosphere is transparent to the incoming solar radiation and in low temperatures during the night (Tmin) because less outgoing longwave radiation is absorbed by the atmosphere. On the other hand, in overcast conditions, Tmax is relatively smaller because a significant part of the incoming solar radiation never reaches the earth's surface and is absorbed and reflected by the clouds. Similarly, Tmin will be relatively higher as the cloud cover acts as a blanket and decreases the net outgoing longwave radiation. Therefore, the difference between the maximum and minimum air temperature (Tmax - Tmin) can be used as an indicator of the fraction of extraterrestrial radiation that reaches the earth's surface. This principle has been utilized by Hargreaves and Samani to develop estimates of ETo using only air temperature data.

**Usage**

estimate\_Rs\_from\_airTemDiff(Ra, Tmax, Tmin, locations)

**Arguments**

Ra	extraterrestrial radiation [MJ m <sup>-2</sup> d <sup>-1</sup> ].
Tmax	maximum air temperature.
Tmin	minimum air temperature.
locations	The adjustment coefficient kRs is empirical and differs for interior' or 'coastal' regions.

**Value**

A vector for solar radiation

**Note**

The temperature difference method is recommended for locations where it is not appropriate to import radiation data from a regional station, either because homogeneous climate conditions do not occur, or because data for the region are lacking. For island conditions, the methodology of Equation 50 is not appropriate due to moderating effects of the surrounding water body. Caution is required when daily computations of ETo are needed. The advice given for Equation 49 fully applies. It is recommended that daily estimates of ETo that are based on estimated Rs be summed or averaged over a several-day period, such as a week, decade or month to reduce prediction error.

**References**

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

---

FIalfalfa	<i>A example dataset of alfalfa under flood irrigation</i>
-----------	--

---

**Description**

A example dataset of alfalfa under flood irrigation

**Usage**

FIalfalfa

**Format**

A data frame with 161 rows and 22 variables

---

Kcb_adj_for_DualKc	<i>Adjust the recommended Kc values at the middle and late stages</i>
--------------------	---

---

**Description**

Adjust the recommended Kc values at the middle and late stages

**Usage**

Kcb\_adj\_for\_DualKc(Kcb\_table, u2, RHmin, h)

**Arguments**

Kcb_table	Recommended value of KC in FAO 56 at the middle and late stages
u2	wind speed at 2 m
RHmin	Minimum relative humidity
h	Plant height

**Value**

A value for adjust Kc at middle and late stages

**References**

Allen, R. G., Pereira, L. S., Raes, D., & Smith, M. FAO Irrigation and drainage paper No. 56. Rome: Food and Agriculture Organization of the United Nations, 1998.

---

`linear_interpolation` *linear interpolation for vector*

---

**Description**

Linear interpolation is performed by using the values on both sides of the missing values.

**Usage**

```
linear_interpolation(DataVector)
```

**Arguments**

DataVector data vector. Note that the starting value of vector needs to be no missing value.

**Value**

A interpolated vector

---

 Model\_DualKc

*Simulation of evapotranspiration using dual crop coefficient method*


---

**Description**

Simulation of evapotranspiration using dual crop coefficient method

**Usage**

```
Model_DualKc(data, param)
```

**Arguments**

data	A data box. Contains the daily data required by the model. You can refer to the function <code>create_modelData()</code>
param	A list. Contains additional parameters. <code>list(Kini,Kmid,Kend,fw,rootDepth,Dei_start,Dri_start,FCe,WPe,Z</code>

**Value**

A list for the model result including a data frame of daily model result ,a list of plots, A data frame of summary data

**Note**

The stages of data should include all four stages. If a crop has multiple growth cycles, each cycle should include all four stages.

**Examples**

```
library(simET)
data("FIalfalfa")
names(FIalfalfa)
#--Model parameter
Dparam_FI<-list(Kini=0.3,#Kcb for initial stage
                Kmid=1.15,#Kcb for mid-season stage
                Kend=1.1,#Kcb for late season stage
                DI=FALSE,#Is it drip irrigation?
                fw=1,# The fraction of the surface wetted
                rootDepth=1.2,#Maximum root depth
                Dei_start=0,#Initial depletion of evaporation layer
                Dri_start=35,#Initial depletion of root layer
                FCe=0.22,#Field capacity of evaporation layer
                WPe=0.15,#Wilting point of evaporation layer
                Ze=0.15,#Depth of the surface soil layer
                REW=6,#Readily evaporable water
                TAW=297,#Total available soil water of the root zone
                p=0.55,#Evapotranspiration depletion factor
                FCrmm=430,#Field capacity of root layer
                CR_param=c(430,-0.32,310,-0.16,-1.4,6.8,1.11,-0.98))
```

```

        #Capillary rise model parameters
    )
  #--Run model
  Model_re_FI<-Model_DualKc(data = FIalfalfa,param = Dparam_FI)
  #--The Result data
  Model_re_FI$Result
  Model_re_FI$Plot
  #--The goodness Of Fit
  estimate_goodnessOfFit(Sim = Model_re_FI$Result$Sim_SoilWater,
                        Obs = Model_re_FI$Result$SoilWater)

```

---

Model_single_Kc	<i>Simulation for evapotranspiration using single crop coefficient method</i>
-----------------	---

---

## Description

Simulation for evapotranspiration using single crop coefficient method

## Usage

```
Model_single_Kc(data, param)
```

## Arguments

data	A data box. Contains the daily data required by the model. You can refer to the function <code>create_modelData()</code>
param	A list. Contains additional parameters.

## Value

A list for the model result including a data frame of daily model result ,a list of plots, A data frame of summary data

## Note

The stages of data should include all four stages. If a crop has multiple growth cycles, each cycle should include all four stages.

## Examples

```

library(simET)
#--Data preparation
data("FIalfalfa")
#--Parameter preparation
param_SingleKc<-list(Kc_mid=1.2,#Kcb for mid-season stage
                    Kc_end=1.15,#Kcb for late season stage
                    rootDepth=1.2,#Maximum root depth
                    #The soil type used for calculating
                    #Kc for initial stage

```

```

soil_type="coarse soil textures",
Dr_start=40,#Initial depletion of root layer
TAW=290,#Total available soil water of the root zone
p=0.55,#Evapotranspiration depletion factor
Field_capacity=420,#Field capacity of root layer
fw=1,#The fraction of the surface wetted
#Capillary rise model parameters
CR_param=c(420,-0.32,303,-0.16,-1.4,6.8,1.11,-0.98)
)

#--Run model
Re_SingleKc<- Model_single_Kc(data = FIalfalfa, param = param_SingleKc)
#--The Result data
Re_SingleKc$Result
Re_SingleKc$Plot
#--The goodness Of Fit
estimate_goodnessOfFit(Sim = Re_SingleKc$Result$Sim_SoilWater,
                        Obs = Re_SingleKc$Result$SoilWater)

```

---

Model\_SW

---

*Simulation of evapotranspiration using Shuttleworth-Wallace model*


---

## Description

Simulation of evapotranspiration using Shuttleworth-Wallace model

## Usage

Model\_SW(data, param)

## Arguments

data	A data box. Contains the daily data required by the model. You can refer to the function <code>create_modelData()</code>
param	A list. Contains additional parameters.

## Value

A list for the model result including a data frame of daily model result ,a list of plots, A data frame of summary data

## Note

The stages of data should include all four stages. If a crop has multiple growth cycles, each cycle should include all four stages.

**Examples**

```

library(simET)
#--Data preparation
data("FIalfalfa")
#--Parameter preparation
param_SW<-list(
  plant=list(
    #the canopy extinction coefficient for net radiation
    kRn=0.3,
    alpha_plant=0.3,#Canopy reflectance
    w=0.01,#Leaf width
    Lmax=10,#Maximum leaf area index
    a1=10,# Leaf stomatal resistance coefficients
    a2=0.005,# Leaf stomatal resistance coefficients
    p=0.5,#the param of reduction factor for T
    rootDepth=1.2 #Maximum root depth
  ),
  Soil=list(
    zs0=0.04,#The soil surface roughnesslength (m)
    tau2=2,#Soil tortuosity
    PHI_p=2, #Soil porosity
    #The soil pore-size distribution index
    #from the Brooks-Corey equation.
    lambda_p=0.18,
    l1=0.02,#Depth of the surface soil layer (m)
    l2=1.2,#Depth of the root layer (m)
    #Saturation water content of evaporation layer
    THETA_v_sat_1=0.36,
    THETA_v_sat_2=0.40, #Saturation water content of root layer
    THETA_start_1=0.2,#Initial water content of evaporation layer
    THETA_start_2=0.36,#Initial water content of root layer
    THETA_wp1=0.15,#Wilting point of evaporation layer
    THETA_wp2=0.15,#Wilting point of root layer
    #Empirical coefficient of evaporation layer.
    #13 for homogenous soil
    alpha1=14,
    alpha2=14,#Empirical coefficient of root layer.
    #Saturated hydraulic conductitivity of evaporation layer
    Ksat_1=13.52,
    Ksat_2=0.02,#Saturated hydraulic conductitivity of root layer
    #Capillary rise model parameters
    CR_param=c(430,-0.32,313,-0.16,-1.4,6.8,0.5,-0.98)
  ),
  Mete=list(
    nu=2,#The wind speed extinction coefficient
    nK=2, #The eddy diffusivity extinction coefficient(taken as 2)
    zr=2, #The reference height (m)
    ##The vapor diffusion coefficient in air (24.7 10-6 m2 s-1)
    Dm_v=24.7*10-6,
    deltaT=1 #Time step difference (day)
  )
)

```

```
#--Run model
Re_SW<-Model_SW(data = FIalfalfa, param = param_SW)
#--The Result data
Re_SW$Result
Re_SW$Plot
#--The goodness Of Fit
estimate_goodnessOfFit(Sim = Re_SW$Result$Sim_SoilWater,
                       Obs = Re_SW$Result$SoilWater)
```

---

SDIalfalfa

*A example dataset of alfalfa under subsurface drip irrigation*

---

### **Description**

A example dataset of alfalfa under subsurface drip irrigation

### **Usage**

SDIalfalfa

### **Format**

A data frame with 161 rows and 22 variables



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